

Testing possible relationships between *Acropora digitifera* genes, seawater chemistry and skeletal elements

TOMOKO BELL,^{1,2*} AKIRA IGUCHI,³ ATSUSHI SUZUKI,⁴ ARISA SEKI² and YUSUKE YOKOYAMA^{1,2}

¹Atmosphere and Ocean Research Institute, The University of Tokyo,
5-1-5, Kashiwanoha, Kashiwa-shi, Chiba 277-8564, Japan

²Department of Earth and Planetary Science, Graduate School of Science, The University of Tokyo,
7-3-1, Hongo, Bunkyo-ku, Tokyo 113-0033, Japan

³Department of Bioresources Engineering, National Institute of Technology, Okinawa College,
Henoko 905, Nago-shi, Okinawa 905-2192, Japan

⁴Geological Survey of Japan, National Institute of Advanced Industrial Science and Technology,
1-1-1, Higashi, Tsukuba, Ibaraki 305-8567, Japan

(Received December 9, 2016; Accepted December 1, 2017)

Coral skeletons are robust tools for examining past environments. However, biogenic effects during skeletal formation cause uncertainties in paleoclimate reconstructions. Thus establishing a method to separate biogenic effects from abiogenic ones during skeletal formation is required. Here we utilized an open access and searchable gene database for the staghorn coral *Acropora digitifera* and examined the number of genes related to the elements in seawater to assess the origin of uncertainties in geochemical proxies. We found that *A. digitifera* has genes that can process at least 15 chemical elements as individual substances (Ca, Na, Zn, K, C, N, Cl, S, Fe, Mg, Mn, Cu, H, Mo, and Te) and transporters for seven of these elements (Ca, Na, Zn, K, Cl, Cu, and H). The number of Ca-related genes was the highest (at least 428 genes, including 53 transporters), whereas Sr, one of the most widely used geochemical proxies, was not found in the gene database. Furthermore, we analyzed skeletal samples of *A. digitifera* exhibiting different growth rates; their Sr/Ca ratios showed the lowest variation (1.9%), whereas other proxies (K/Ca, Na/Ca, and Mg/Ca) showed higher variation (2.3–11.9%). This might be linked to the number of genes related to the proxies (namely, the magnitude of biogenic and/or abiogenic effects). We suggest that considering elements with no relevant coral genes could provide effective criteria for reliable proxies (e.g., Sr/Ca, Li/Ca and U/Ca).

Keywords: seawater element, geochemical proxies, genomic information, scleractinian coral, *Acropora digitifera*

INTRODUCTION

Scleractinian corals (including the staghorn corals) form massive amounts of calcium carbonate (CaCO₃), and their calcification rate is estimated to be 2–6 kg CaCO₃ m⁻² yr⁻¹ (Barnes and Devereux, 1984). Coral reefs cover an area of over 284,300 km² globally (Spalding *et al.*, 2001), which is ~0.08% of the entire ocean surface. Given that ocean environments (e.g., seawater elements and physical conditions such as temperature) influence the composition and concentration of elements in coral skeletons, they have been used as effective geochemical tools to investigate past environments (e.g., Yokoyama and Esat, 2015; Yokoyama *et al.*, 2011). For example, Sr/Ca

ratios in coral skeletons have been widely used to reconstruct sea surface temperature (SST) (e.g., Gagan *et al.*, 2000; Seki *et al.*, 2012; Felis *et al.*, 2014), and the calcium accumulation in coral skeletons has been explained as mainly inorganic precipitation (Kinsman and Holland, 1969; Dietzel *et al.*, 2004). Also, a variety of chemical ratios (e.g., U/Ca, Li/Mg, and Cd/Ca), have been applied as geochemical proxies (e.g., Inoue *et al.*, 2011; Montagna *et al.*, 2014; Pretet *et al.*, 2014). One downside to using coral skeletons as geochemical proxies is the potential uncertainties related to biomineralization processes (e.g., biogenic effects) such as growth-rate-related kinetic effects (e.g., Hayashi *et al.*, 2013; Hirabayashi *et al.*, 2013) and individual variability (e.g., Bell *et al.*, 2017). If the magnitude of these effects during skeletal formation is high, coral skeletons cannot be considered robust tools for reconstructing past environments. However, the mechanisms causing these uncertainties in coral skeletal proxies remain unclear. For example, Mg/Ca ratios have

*Corresponding author (e-mail: tbell@aori.u-tokyo.ac.jp)

been well studied as SST proxies, but most of the studies reported lower correlations between SST and Sr/Ca ratios (e.g., Inoue *et al.*, 2007). To elucidate this phenomenon, Wang and Xu (2001) suggested an energy partition model to explain the difference between Sr and Mg. They calculated that the distribution coefficients of Sr, Mg, and Ca were 1.26, 0.0087, and 1, respectively, in aragonite. They suggested that Sr has stronger ability to bind to Ca abiogenically than Mg does. Gene composition from genome information, as related to specific elements, could help us understand this aspect. Nevertheless, there are no systematic explanations by which to evaluate the biogenic effects observed in geochemical proxies.

Owing to recent progress in providing coral genome information through increased throughput of next-generation sequencing technologies (Meyer *et al.*, 2009; Shinzato *et al.*, 2011), molecular-based research on biological processes of corals has also advanced rapidly (e.g., biomineralization; Moya *et al.*, 2012; Vidal-Dupiol *et al.*, 2013). Several molecules have been proposed to form the molecular basis of coral calcification (e.g., Ca²⁺ ATPase (Zoccola *et al.*, 2004), carbonic anhydrase (Moya *et al.*, 2008), and organic matrix protein (Watanabe *et al.*, 2003), yet the molecules related to seawater elements (except for Ca) have not been fully explored. The entire genome sequence of *A. digitifera*, one of the dominant species in the coral reefs of Japan, has been decoded (Shinzato *et al.*, 2011) and data on the gene components of the species are also available (Dunlap *et al.*, 2013). Thus, this species is an ideal model for studying and understanding the molecular basis of coral calcification.

In this study, to elucidate the factors causing uncertainties in skeletal proxies, we searched for genes related to elemental metabolism using the genomic information of *A. digitifera*. We pursued the possibility that the existence of these genes could be used to distinguish biological influences on the formation of coral skeletons. Thus, we attempted to identify molecules specialized for metabolizing individual elements because this database allows us to search proteins by inputting elemental terms. For ions in seawater to be moved into coral bodies biogenically, corals need to have specific genes to transport them. Therefore, we hypothesized that the number and type of genes related to each element would be correlated with the magnitude of the relevant biological functions, and this might contribute to uncertainties on skeletal proxies affected by biological processes. In addition, we analyzed skeletal samples of *A. digitifera* exhibiting a wide range of growth rates in a common garden culture experiment (described below) to confirm the elemental variation in skeletal proxies. To our knowledge, this is the first study to combine studies of coral genomes and geochemistry to elucidate the robustness of geochemical proxies.

MATERIALS AND METHODS

Inventory of metabolic and transporter genes for seawater elements

To identify the genes in the coral genome related to metabolizing seawater elements, we used published data from ZoophyteBase (Dunlap *et al.*, 2013). This allowed us to search for genes of *A. digitifera* that were predicted to produce the functional proteins that perform metabolic processes involving seawater elements. In ZoophyteBase, proteins of *A. digitifera* are annotated using hidden Markov models and the Kyoto Encyclopedia of Genes and Genomes (KEGG). The latter is widely used as a resource for investigating biological systems (Dunlap *et al.*, 2013) and for linking genes to higher-level functions. With this search engine, we used chemical elements found in seawater as keywords to search for functional proteins related to each element. Although not all of the elements in the periodic table are present in seawater, we chose to investigate every element experimentally. In general, the simple substances in coral skeletons are analyzed as temperature proxies, and chemical compounds are not used. Therefore, it should be noted that chemical compounds, such as bicarbonate, were not investigated in our analysis although bicarbonate transporters in coral (*Acropora millepora*) have been studied (Moya *et al.*, 2012). When a gene was found to have multiple local sequences expressing different functional proteins, that gene was counted based on the number of functional proteins. For example, the gene *adi_v1.09241* appeared three times when the search for “calcium” was conducted because this gene consisted of three partial amino acid sequences that could express three different functional proteins. In this case, *adi_v1.09241* was counted as three genes in this study. In addition, the number of transporters was counted for each element from ZoophyteBase. This is because transporters are known to regulate ion flows in various types of eukaryotic cell membranes and might have a functional role in calcification. Transporters control the movement of ions across cell membranes by binding to target ions, and if present in calcicoblastic cell membranes, they control ion flows to extracytoplasmic calcifying fluid (ECF), where calcification occurs. According to Lodish *et al.* (2000), transporters can carry ~10⁴ ions per second; thus, the inventory of transporter genes is informative for estimating the range and potentially the amount of ion transport from seawater to ECF.

ICP-AES analysis of coral skeletons

Several colonies of *A. digitifera* were selected from a fringing reef of Sesoko Island (Okinawa, Japan) in 2012. The branches of *A. digitifera* collected from these colonies were at least 3 m apart. All the samples were collected in strict accordance with good animal practice de-

Table 2. Number of metabolic and transporter genes

Chemical element	Number of metabolic genes	Number of transporters
Ca	428	53
Na	173	76
Zn	149	20
K	131	39
C	127	0
N	54	0
Cl	44	24
S	43	0
Cu	7	4
H	6	6
Fe	3	0
Mo	3	0
Mg	1	0
Mn	1	0
Te	1	0

(from 0 to 428 hits), and the chemical elements were aligned in order according to these numbers (Table 2). Because we were interested in these elements as geochemical tools, the elements that have been previously used as temperature proxies are marked with asterisks in Table 1. There are 90 elements that occur in nature, and 15 of these had hits in the ZoophyteBase search (Ca, Na, Zn, K, C, N, Cl, S, Fe, Mg, Mn, Cu, H, Mo, and Te). Therefore, *A. digitifera* has metabolic genes to process at least these 15 seawater elements as simple substances. Among the 15 seawater elements that could be processed by the metabolic genes of *A. digitifera*, five elements were major seawater elements whose concentrations are >1 ppm in seawater (Cl, Na, Mg, Ca, and K), and seven elements (Zn, S, Fe, Mn, Cu, Mo, and Te) were trace elements that are defined as those with concentrations of <1 ppm. According to Pilson (1998), there are eight major elements in seawater that occur as simple substances: Cl (545.120 mmol/kg), Na (469.00 mmol/kg), Mg (52.82 mmol/kg), Ca (10.27 mmol/kg), K (10.21121 mmol/kg), Br (0.842 mmol/kg), Sr (0.090 mmol/kg), and F (0.068 mmol/kg). Of these eight major seawater elements, four were within the top 10 elements with the highest number of genes in Table 2: Ca (428 genes), Na (173), K (131), and Cl (44). Interestingly, genes related to three other major elements in seawater (Sr, Br, and F) were not found in the ZoophyteBase search. Thus, it is possible that there are no specific metabolic genes to process these elements in spite of their high concentrations in seawater. It was noteworthy that Sr, which is a major seawater element and the element most used as a geochemical proxy, was not found in the database. Although Ca had only the fourth highest concentration (411.9 ppm) among the five major seawater elements, the number of genes related to Ca was the highest (428 genes). Moreover, the number of Ca-related metabolic genes was more than two times greater

than the number of genes for the second-ranking element, Na (173 genes).

The appearance of new genes and earth environments might be partly correlated. For example, the appearance of manganese catalase, which breaks down hydrogen peroxide into oxygen is tied to a major oxygenation event about 2.4 billion years ago (Kim *et al.*, 2012). In this study, we compared the concentration of each element in seawater and the number of corresponding genes to see if we could find any correlations between the number of metabolic genes and the elemental concentrations in seawater (Fig. 1). No significant correlations were confirmed in the following three types of groups: (1) major and trace seawater elements together ($p = 0.16$, $r = 0.46$), (2) only major elements ($p = 0.41$, $r = -0.48$), and (3) only trace elements ($p = 0.64$, $r = 0.24$). This indicates that, at least using our approach, the numbers of metabolic genes are not determined simply by the chemical composition and concentration of the elements in seawater. It should be noted that the numbers of genes related to each element are associated with not only coral calcification but also other metabolic processes. For example, calmodulin, which is one of the Ca^{2+} -sensing proteins and was found in the coral genome (Dunlap *et al.*, 2013), has many regulatory roles such as cytoskeleton organization, cellular metabolism, and cell differentiation (Benaim and Villalobo, 2002; Cyert, 2001; Vetter and Leclerc, 2003). Thus, calcium is involved in many metabolic processes in addition to skeletal formation in the corals mentioned above.

Transporter genes and seawater chemistry

Of the 15 elements that are related to the metabolic genes of *A. digitifera*, we found that seven elements had related transporters (Na, K, Cl, Zn, Ca, Cu, and H; Table 2). Although we cannot identify where the transporters

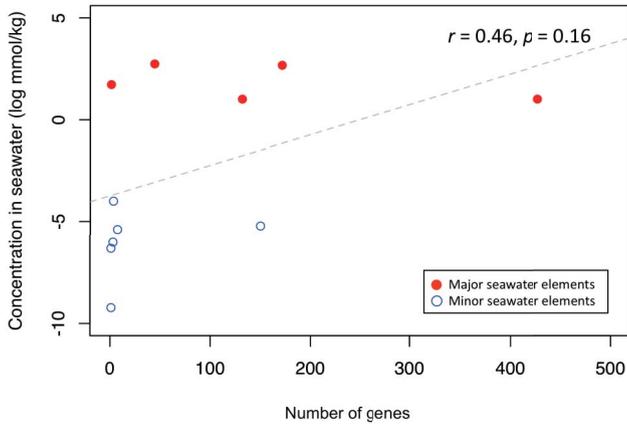


Fig. 1. Cross-plots of the number of genes and respective seawater concentrations of Ca, Na, K, Cl, and Mg (major seawater elements, red solid circles); and Zn, Mn, Cu, Fe, Te, and Mo (minor seawater elements, blue circles). The seawater concentrations are from Pilson (1998).

are located, this finding suggests that at least these seven elements obtained from seawater are physically passed across membranes by transporters, possibly to reach the ECF of *A. digitifera*. In the ZoophyteBase search, the most relevant transporters were found for Na (76 transporters), Ca (53), and K (39). These three elements are all major seawater elements, and the transporters related to these elements may import or export these elements between cells and seawater. Interestingly, the trace elements Zn and Cu, which are heavy metals known to have important biological roles (Morel and Price, 2003) as cofactors of many enzymes (Bisc er e *et al.*, 2015), were among the seven elements for which relevant transporters were found. We also searched for correlations between the numbers of transporter genes and mineral concentrations in seawater (Fig. 2). A significant correlation was confirmed in the first of the following three types of groups: (1) major and trace seawater elements ($p = 0.02$, $r = 0.68$), (2) only major elements ($p = 0.87$, $r = 0.10$), and (3) only trace elements ($p = 0.61$, $r = 0.26$). This indicates that the number of transporters could possibly be influenced by seawater chemistry; however, we need to collect more data from other taxa (e.g., other cnidarians) to confirm this trend.

ICP-AES analysis of *A. digitifera* skeletons

A total of 13 branches from five different colonies of *A. digitifera* showed different growth rates (0.23–1.61%/d) after the common-garden culture experiment (Fig. 3). The growth rates were significantly different among colonies ($F(4, 8) = 23.898$; $p < 0.001$). Because it is a regular protocol in geochemical applications to normalize the skeletal element concentrations using the Ca concentra-

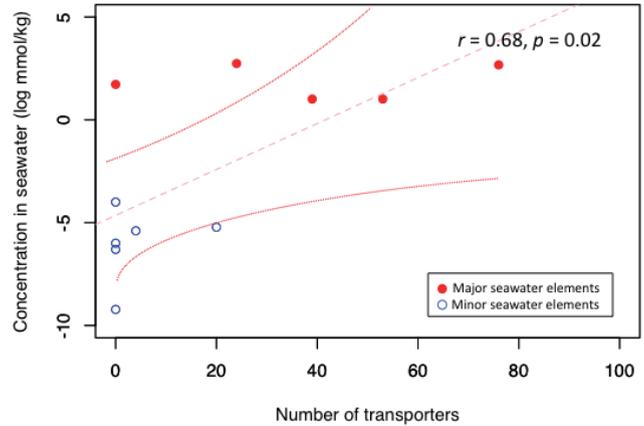


Fig. 2. Cross-plots of the number of transporters and respective seawater concentrations of Ca, Na, K, Cl, and Mg (major seawater elements, red solid circles); and Zn, Mn, Cu, Fe, Te, and Mo (minor seawater elements, blue circles). Significant correlation was observed between the number of transporters and seawater concentration ($p = 0.02$, $r = 0.68$), and the pink dotted line represents the regression line: $y = 0.112x - 4.646$. The red dotted lines represent the 95% confidence interval.

tion, we calculated Sr/Ca, K/Ca, Na/Ca, and Mg/Ca for these 13 samples (Fig. 4). The Sr/Ca, Na/Ca, and Mg/Ca values are within the ranges previously reported by Mitsuguchi *et al.* (2010), Gagan *et al.* (2000) and Inoue *et al.* (2007). Our K/Ca values (0.21–0.24 mmol/mol) are similar to those reported by Mitsuguchi and Kawakami (2012) for *Porites* sp. (0.16–0.22 mmol/mol). The values of these proxies were not significantly different among colonies ($F(4, 8) < 3.69$; $p > 0.05$). However, intriguingly, the Mg values (0.072–0.098 mmol) were significantly different among colonies ($F(4, 8) > 5.4329$; $p < 0.05$) while those of Sr (0.14–0.18 mmol) were not ($F(4, 8) < 0.2751$; $p > 0.05$). Thus, without standardization of calcium, magnesium showed colony-specific values. In geochemical studies, only one or two coral core samples are typically collected for skeletal-element analysis because geochemists are required to collect the most samples possible from one coral colony to obtain a few hundred years' worth of data. Thus, it is not feasible or realistic to collect more than three coral core samples. Therefore, to be a strong geochemical tool, a proxy based on coral skeletal elements should be consistent among colonies from the same environment, regardless of variations in growth rates. The standard deviations of Sr/Ca, K/Ca, Na/Ca, and Mg/Ca were (1.9, 4.2, 2.3, and 11.9)%, respectively. Thus, Sr/Ca showed the smallest variation and Mg/Ca showed the largest. Hayashi *et al.* (2013) also reported stable Sr/Ca ratios among *Porites australiensis* samples with a four-fold growth-rate variation, which is a smaller range than found in this study (sevenfold variation). The wide range of Mg/Ca ratios of *Porites* spp. has been reported in many



Fig. 3. Photo of six *A. digitifera* samples from two different colonies that showed the fastest and slowest growth rates (0.23 and 1.61%/d, respectively) after 12 months in an outside aquarium at the Sesoko Station, Tropical Biosphere Research Center, University of the Ryukyus, Japan.

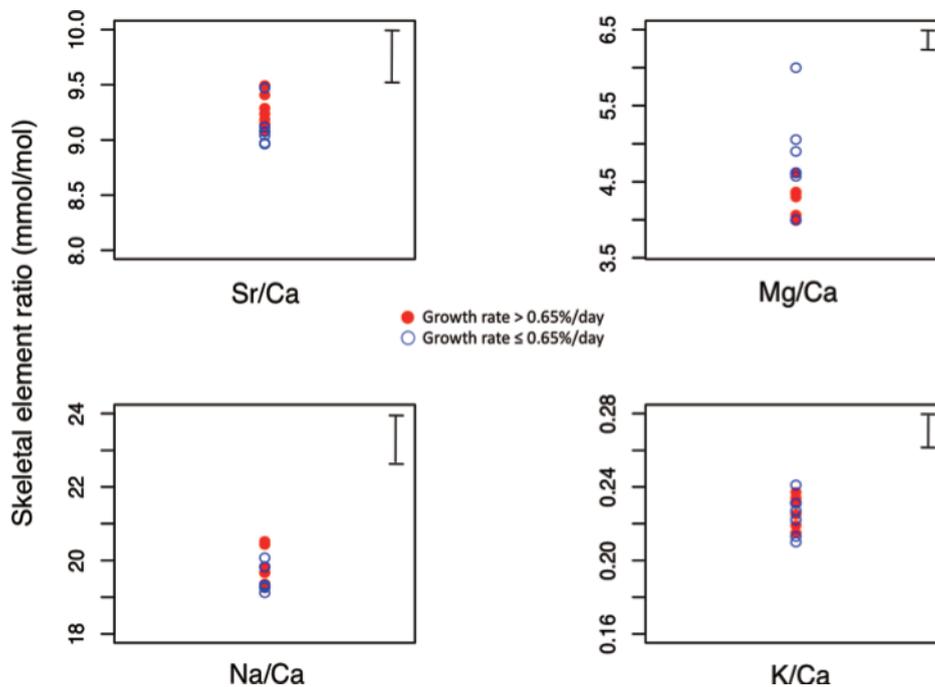


Fig. 4. Skeletal element ratios of 13 *A. digitifera* samples that showed varied growth rates (growth rate > 0.65%/d: red solid circles, growth rate < 0.65%/d: blue circles). The values of these proxies were not significantly different among colonies ($F(4, 8) < 3.69$). The error bar indicates the maximum difference observed between two kinds of standards: XSTC-13 (multi-element standard solution, SPEX) and JCp-1 (a coral standard material, Geological Survey of Japan).

studies (e.g., Inoue *et al.*, 2007), and their growth rates are considered to be related to the Mg/Ca ratios in their skeletons. Therefore, we report the influence of growth rate on skeletal elements next.

Influence of growth rate on skeletal elements

The correlation between skeletal elements (Sr/Ca, K/Ca, Na/Ca, and Mg/Ca) and growth rates were various and are summarized in Table 3. The correlation between growth rate and the Mg/Ca ratio in this study was evident, as previously reported using *Porites* spp., and the correlation could be due to the difference in the partition

coefficient between Mg and Ca in aragonite (Inoue *et al.*, 2007). We also confirmed a significant correlation between growth rate and the Sr/Ca ratio in *A. digitifera*. However, the correlation coefficients between skeletal elements and growth rates were -0.69 ($p < 0.01$) and 0.24 ($p = 0.42$) for Mg and Sr, respectively. Thus, Mg was correlated with the growth rate compared to Sr, and one possible scenario was that Mg was acting as glue to cement molecules of CaCO_3 (Bell *et al.*, 2017). Interestingly, the variation in the Sr/Ca ratio of *A. digitifera* was only 1.9%, which was one-sixth of the variation in the Mg/Ca ratio (11.9%). Thus, the influence of the growth

Table 3. Correlations between growth rates and skeletal elements

	Sr/Ca	Mg/Ca	Na/Ca	K/Ca
Growth rate	0.55 ($p = 0.05$)	-0.70 ($p = 0.01$)	0.19 ($p = 0.53$)	0.11 ($p = 0.72$)
Sr/Ca	—	-0.53 ($p = 0.06$)	0.45 ($p = 0.13$)	0.40 ($p = 0.17$)
Mg/Ca	—	—	-0.27 ($p = 0.37$)	0.23 ($p = 0.44$)
Na/Ca	—	—	—	0.54 ($p = 0.06$)
K/Ca	—	—	—	—

Significant correlations are written in bold letters.

rate on Sr/Ca is much smaller than that on Mg/Ca in *A. digitifera*. Although there are no comparable studies investigating the relationship between growth rate and the K/Ca or Na/Ca ratio, our results show no correlations between these ratios and growth rate. It is noteworthy that Mitsuguchi and Kawakami (2012) reported a positive correlation between K/Ca and Na/Ca, and our result also showed a positive correlation ($r = 0.54$, $p = 0.06$; Table 3). However, we need more data to compare our study to theirs because the number of our samples ($N = 13$) was many fewer than that of Mitsuguchi and Kawakami (2012) ($N = 60$).

Relationships between the number of genes, growth rates and skeletal elements

Our results using Zoophybase showed that *A. digitifera* depends on at least 15 chemical elements in terms of genomic information. In general, the increase in the number of genes is related to biological complexity (Alberts *et al.*, 2002). Moreover, Weng and Noel (2012) stated that metabolism is a complex network of chemical transformations mediated by a multitude of enzymes. Therefore, the number of genes can be an index of the complicity of metabolic processes. Considering the larger number of Ca-related genes in the coral genome, the implication is that a stable and well-developed Ca metabolic system exists in corals, and that Ca is metabolized in a regular and steady manner biogenically. However, we need to caution readers that the genes related to Ca are not always involved in the biogenic process of coral calcification (see the section, Metabolic Genes and Seawater Chemistry). Large numbers of genes related to Na and K exist, and they are also controlled biogenically by metabolic processes in *A. digitifera*. On the other hand, the chemical elements that exist in seawater but were not found in the gene database (e.g., Sr, Li, and U) might be processed mainly abiogenically. The Sr/Ca values had a variation of only 1.9%, regardless of the growth rate of the coral colony, and the Mg/Ca values showed the largest variation (11.9%). We suggest that elements with no relevant coral genes could be good candidates for proxies with fewer biogenic effects (e.g., Sr/Ca, Li/Ca, and U/Ca). For example, Li/Ca and U/Ca were proposed as

useful temperature proxies (e.g., Hathorne *et al.*, 2013; Min *et al.*, 1995); moreover, DeCarlo *et al.* (2016) recently introduced Sr-U as a reliable temperature proxy by combining Sr/Ca and U/Ca. Thus, the number of genes in coral genomes related to specific elements may provide at least partial criteria for determining reliable proxies.

At this stage, from the viewpoint of molecular physiology, it is unclear why the high correlation between Mg/Ca and growth rate was observed or why the variation in Mg/Ca was the highest. Tanaka *et al.* (2015) conducted *A. digitifera* culture experiments at three different pH levels and reported that the variation of Mg/Ca among pH conditions was statistically insignificant. Moreover, the variation of Mg/Ca among three colonies was wider than that for two other colonies (4.16–5.21 mmol/mol). This implies strong biological control over skeletal Mg/Ca during abiogenic fluctuations (e.g., pH). They suggested that incorporation of Mg/Ca is to some extent also affected by physiological, colony-specific factors. These uncertainties regarding Mg/Ca could possibly be explained by our ongoing study focusing on Mg-related gene expression among samples with different growth rates. Another intriguing finding in this study, the possible correlation between the number of transporters and specific elements in seawater should be investigated using other non-calcifying cnidarians such as *Nematostella vectensis* and *Hydra magnipapillata*, because their entire genome sequences are also available (Putnam *et al.*, 2007; Chapman *et al.*, 2010). On the other hand, it should be noted that results focusing on the number of genes need to be interpreted with some caution because there are some unknown genes in the coral genome (Shinzato *et al.*, 2011) that might contribute to the elemental composition of coral skeletons. Thus, comprehensive gene expression profiling (such as RNA-seq) in the future, using coral branches with different growth rates, would provide more details on biogenic effects during skeletal formation. In addition, it is necessary to understand how ambient seawater is connected to calcification sites (ECF) in corals, namely, whether there are transcellular or paracellular pathways of skeletal materials to the ECF sites (e.g., Gagnon *et al.*, 2012; Ohno *et al.*, 2017).

CONCLUSIONS

One of the most important topics in biomineralization research is whether chemical elements in marine calcifiers are controlled biogenically or abiogenically when they are transported from seawater into the skeleton. To resolve this issue, we propose that a bioinformatics approach using genome information would be an effective method. We suggest that elements with no relevant coral genes would make good candidates for reliable proxies. Genomic information can help us find new geochemical proxies with the fewest biogenic effects, and also explain the robustness of proxies that are already known to be effective in reconstructing past ocean environments.

Acknowledgments—The data used in this manuscript are available upon request to readers. This research was partly supported by the CANON Foundation (R12-Z-0013) and JSPS KAKENHI (26247085). We thank the Japanese Municipality of Okinawa Prefecture for the permit to collect the coral samples, and M. Inoue (Okayama University) and K. Kubota (The University of Tokyo) for their valuable advice on the ICP-AES analysis.

REFERENCES

- Alberts, B., Johnson, A., Lewis, J., Raff, M., Roberts, H. and Walter, P. (2002) *Molecular Biology of the Cell*. 4th ed., Garland Science, Chap. 7.
- Barnes, D. J. and Devereux, M. J. (1984) Productivity and calcification on a coral reef: a survey using pH and oxygen electrode techniques. *J. Exp. Mar. Biol. Ecol.* **9**, 213–231.
- Bell, T., Nishida, K., Ishikawa, K., Suzuki, A., Nakamura, T., Sakai, K., Iguchi, A. and Yokoyama, Y. (2017) Temperature controlled culture experiments using primary polyps of coral *Acropora digitifera*: their calcification rate variations and skeletal Sr/Ca, Mg/Ca, and Na/Ca ratios. *Paleogeogr. Paleoclimatol. Paleocol.* **484**, 129–135.
- Benaïm, G. and Villalobo, A. (2002) Phosphorylation of calmodulin. Functional implications. *Eur. J. Biochem.* **269**, 3619–3631.
- Biscéré, T., Rodolfo-Metalpa, R., Lorrain, A., Chauvaud, L., Thébault, J., Clavier, J. and Houlbrèque, T. (2015) Responses of two Scleractinian corals to cobalt pollution and ocean acidification. *PLoS ONE* **10**, e0122898, doi:10.1371/journal.pone.0122898.
- Chapman, J. A., Kirkness, E. F., Simakov, O., Hampson, S. E., Mitros, T., Weinmaier, T., Rattei, T., Balasubramanian, P. G., Borman, J., Busam, D., Disbennett, K., Pfannkoch, C., Sumin, N., Sutton, G. G., Viswanathan, L. D., Walenz, B., Goodstein, D. M., Hellsten, U., Kawashima, T., Prochnik, S. E., Putnam, N. H., Shu, S., Blumberg, B., Dana, C. E., Gee, L., Kibler, D. F., Law, L., Lindgens, D., Martinez, D. E., Peng, J., Wigge, P. A., Bertulat, B., Guder, C., Nakamura, Y., Ozbek, S., Watanabe, H., Khalturin, K., Hemmrich, G., Franke, A., Augustin, R., Fraune, S., Hayakawa, E., Hayakawa, S., Hirose, M., Hwang, J. S., Ikeo, K., Fujisawa, C. N., Ogura, A., Takahashi, T., Steinmetz, P. R. H., Zhang, X., Aufschnaiter, R., Eder, M. K., Gorný, A. K., Salvenmoser, W., Heimberg, A. M., Wheeler, B. M., Peterson, K. J., Bottger, A., Tischler, P., Wolf, A., Gojbori, T., Remington, K. A., Strausberg, R. L., Craig Venter, J., Technau, U., Hobmayer, B., Bosch, T. C. G., Holstein, T. W., Fujisawa, T., Bode, H. R., David, C. N., Rokhsar, D. S. and Steele, R. E. (2010) The dynamic genome of *Hydra*. *Nature* **464**, 592–596.
- Cyert, M. S. (2001) Genetic analysis of calmodulin and its targets in *Saccharomyces cerevisiae*. *Annu. Rev. Genet.* **35**, 647–672.
- Davies, P. S. (1989) Short-term growth measurements of corals using an accurate buoyant weighing technique. *Mar. Biol.* **101**, 389–395.
- DeCarlo, T. M., Gaetani, G. A., Cohen, A. L., Foster, G. L., Alpert, A. E. and Stewart, J. A. (2016) Coral Sr-U thermometry. *Paleoceanography* **31**, 626–638.
- Dietzel, M., Gussone, N. and Eisenhauer, A. (2004) Co-precipitation of Sr²⁺ and Ba²⁺ with aragonite by membrane diffusion of CO₂ between 10 and 50°C. *Chem. Geol.* **203**, 139–151.
- Dunlap, W. C., Starcevic, A., Baranasic, D., Diminic, J., Zucko, J., Gacesa, R., van Oppen, M. J. H., Hranueli, D., Cullum, J. and Long, P. F. (2013) KEGG orthology-based annotation of the predicted proteome of *Acropora digitifera*: ZoophyteBase—an open access and searchable database of a coral genome. *BMC Genomics* **14**, doi:10.1186/1471.
- Felis, T., McGregor, H. V., Linsley, B. K., Tudhope, A. W., Gagan, M. K., Suzuki, A., Inoue, M., Thomas, A. L., Esat, T. M., Thompson, W. G., Tiwari, M., Potts, D. C., Mudelsee, M., Yokoyama, Y. and Webster, J. M. (2014) Intensification of the meridional temperature gradient in the Great Barrier Reef following the Last Glacial Maximum. *Nat. Commun.* **5**, doi:10.1038/ncomms5102.
- Gagan, M. K., Ayliffe, L. K., Beck, J. W., Cole, J. E., Druffel, E. R. M., Dunbar, R. B. and Schrag, D. P. (2000) New views of tropical paleoclimates from corals. *Quat. Sci. Rev.* **19**, 45–64.
- Gagnon, A. C., Adkins, J. F. and Erez, J. (2012) Seawater transport during coral biomineralization. *Earth Planet. Sci. Lett.* **329**, 150–161.
- Hathorne, E. C., Felis, T., Suzuki, A., Kawahata, H. and Cabioch, G. (2013) Lithium in the aragonite skeletons of massive *Porites* corals: A new tool to reconstruct tropical sea surface temperatures. *Paleoceanography* **28**, 143–152.
- Hayashi, E., Suzuki, A., Nakamura, T., Iwase, A., Ishimura, T., Iguchi, A., Sakai, K., Okai, T., Inoue, M., Araoka, D., Murayama, S. and Kawahata, H. (2013) Growth-rate influences on coral climate proxies tested by a multiple colony culture experiment. *Earth Planet. Sci. Lett.* **362**, 198–206.
- Hirabayashi, S., Yokoyama, Y., Suzuki, A., Kawakubo, Y., Miyairi, Y., Okai, T. and Nojima, S. (2013) Coral growth-rate insensitive Sr/Ca as a robust temperature recorder at the extreme latitudinal limits of *Porites*. *Geochem. J.* **47**, e1–e5.
- Inoue, M., Suzuki, A., Nohara, M., Hibino, K. and Kawahata, H. (2007) Empirical assessment of coral Sr/Ca and Mg/Ca ratios as climate proxies using colonies grown at different temperatures. *Geophys. Res. Lett.* **34**, L12611.
- Inoue, M., Suwa, R., Suzuki, A., Sakai, K. and Kawahata, H.

- (2011) Effects of seawater pH on growth and skeletal U/Ca ratios of *Acropora digitifera* coral polyps. *Geophys. Res. Lett.* **38**, L12809, doi:10.1029/2011GL047786.
- Kim, K. M., Qin, T., Jiang, Y. Y., Chen, L. L., Xiong, M., Caetano-Anollés, D., Zhang, H. Y. and Caetano-Anollés, G. (2012) Protein domain structure uncovers the origin of aerobic metabolism and the rise of planetary oxygen. *Structure* **20**, 67–76.
- Kinsman, D. J. J. and Holland, H. D. (1969) The co-precipitation of cations with CaCO₃-IV. The co-precipitation of Sr²⁺ with aragonite between 16 and 96°C. *Geochim. Cosmochim. Acta* **33**, 1–17.
- Lodish, H., Berk, A., Zipursky, S. L., Matsudaira, P., Baltimore, D. and Darnell, J. (2000) *Molecular Cell Biology*. W. H. Freeman, Chap. 15.
- Meyer, E., Aglyamova, G. V., Wang, S., Buchanan-Carter, J., Abrego, D., Colbourne, J. K., Willis, B. L. and Matz, M. V. (2009) Sequencing and de novo analysis of a coral larval transcriptome using 454 GSFlx. *BMC Genomics* **10**, doi:10.1186/1471-2164-10-219.
- Min, G. R., Edwards, R. L., Taylor, F. W., Recy, J., Gallup, C. D. and Beck, J. W. (1995) Annual cycles of UCa in coral skeletons and UCa thermometry. *Geochim. Cosmochim. Acta* **59**, 2025–2042.
- Mitsuguchi, T. and Kawakami, T. (2012) Potassium and other minor elements in *Porites* corals: implications for skeletal geochemistry and paleoenvironmental reconstruction. *Coral Reefs* **31**, 671–681.
- Mitsuguchi, T., Uchida, T. and Matsumoto, E. (2010) Na/Ca variability in coral skeletons. *Geochem. J.* **44**, 261–273.
- Montagna, P., McCulloch, M., Douville, E., López Correa, M., Trotter, J., Rodolfo-Metalpa, R., Dissard, D., Ferrier-Pagès, C., Frank, N., Freiwald, A., Goldstein, S., Mazzoli, C., Reynaud, S., Rüggeberg, A., Russo, S. and Taviani, M. (2014) Li/Mg systematics in scleractinian corals: Calibration of the thermometer. *Geochim. Cosmochim. Acta* **132**, 288–310.
- Morel, F. M. M. and Price, N. M. (2003) The biogeochemical cycles of trace metals in the Oceans. *Science* **300**, 944–947.
- Moya, A., Tambutté, S., Bertucci, A., Tambutté, E., Lotto, S., Vullo, D., Supuran, C. T., Allemand, D. and Zoccola, D. (2008) Carbonic anhydrase in the scleractinian coral *Stylophora pistillata* characterization, localization, and role in biomineralization. *J. Biol. Chem.* **283**, 25475–25484.
- Moya, A., Huisman, L., Ball, E. E., Hayward, D. C., Grasso, L. C., Chua, C., Woo, H. N., Gattuso, J. P., Foret, S. and Miller, D. J. (2012) Whole transcriptome analysis of the coral *Acropora millepora* reveals complex responses to CO₂-driven acidification during the initiation of calcification. *Mol. Ecol.* **21**, 2440–2454.
- Ohno, Y., Iguchi, A., Shinzato, C., Inoue, M., Suzuki, A., Sakai, K. and Nakamura, T. (2017) An aposymbiotic primary coral polyp counteracts acidification by active pH regulation. *Sci. Rep.* **7**, 40324.
- Okai, T., Suzuki, A., Kawahata, H., Terashima, S. and Imai, N. (2002) Preparation of a new Geological Survey of Japan geochemical reference material: Coral JCp-1. *Geostand. Newsl.* **26**, 95–99.
- Pilson, M. E. Q. (1998) *An Introduction to the Chemistry of the Sea*. Prentice Hall, Chap. 4.
- Pretet, C., Reynaud, S., Ferrier-Pagès, C., Gattuso, J., Kamber, B. S. and Samankassou, E. (2014) Effect of salinity on the skeletal chemistry of cultured scleractinian zooxanthellate corals: Cd/Ca ratio as a potential proxy for salinity reconstruction. *Coral Reefs* **33**, 169–180.
- Putnam, N. H., Srivastava, M., Hellsten, U., Dirks, B., Chapman, J., Salamov, A., Terry, A., Shapiro, H., Lindquist, E., Kapitonov, V. V., Jurka, J., Genikhovich, G., Grigoriev, I. V., Lucas, S. M., Steele, R. E., Finnerty, J. R., Technau, U., Martindale, M. Q. and Rokhsar, D. S. (2007) Sea anemone genome reveals ancestral eumetazoan gene repertoire and genomic organization. *Science* **317**, 86–94.
- R Development Core Team (2013) *R: A language and environment for statistical computing*. R Foundation for Statistical Computing, Vienna, Austria.
- Seki, A., Yokoyama, Y., Suzuki, A., Kawakubo, Y., Okai, T., Miyairi, Y., Matsuzaki, H., Namizaki, N. and Kan, H. (2012) Mid-Holocene sea-surface temperature reconstruction using fossil corals from Kume Island, Ryukyus, Japan. *Geochem. J.* **46**, e27–e32.
- Shinzato, C., Eiichi, S., Kawashima, T., Hamada, M., Kanako, H., Makiko, T., Fujie, M., Fujiwara, M., Koyanagi, R., Ikuta, T., Fujiyama, A., Miller, D. J. and Satoh, N. (2011) Using the *Acropora digitifera* genome to understand coral responses to environmental change. *Nature* **476**, 320–323.
- Spalding, M. D., Ravilious, C. and Green, E. P. (2001). *World Atlas of Coral Reefs*. University of California Press, 10 pp.
- Tanaka, K., Holcomb, M., Takahashi, A., Kurihara, H., Asami, R., Shinjo, R., Sowa, K., Rnkenburg, K., Watanabe, T. and McCulloch, M. (2015) Response of *Acropora digitifera* to ocean acidification: constraints from δ¹¹B, Sr, Mg, and Ba compositions of aragonitic skeletons cultured under variable seawater pH. *Coral Reefs* **34**, 1139–1149.
- Vetter, S. W. and Leclerc, E. (2003) Novel aspects of calmodulin target recognition and activation. *Eur. J. Biochem.* **270**, 404–414.
- Vidal-Dupiol, J., Zoccola, D., Tambutté, E., Grunau, C., Cosseau, C., Smith, K. M., Freitag, M., Dheilily, N. M., Allemand, D. and Tambutté, S. (2013) Genes related to ion-transport and energy production are upregulated in response to CO₂-Driven pH decrease in corals: new insights from transcriptome analysis. *PLoS ONE* **8**, e58652.
- Wang, Y. F. and Xu, H. F. (2001) Prediction of trace metal partitioning between minerals and aqueous solutions: A linear free energy correlation approach. *Geochim. Cosmochim. Acta* **65**, 1529–1543.
- Watanabe, T., Fukuda, I., China, K. and Isa, Y. (2003) Molecular analyses of protein components of the organic matrix in the exoskeleton of two scleractinian coral species. *Comp. Biochem. Physiol. Part B: Biochem. Mol. Biol.* **136**, 767–774.
- Weng, J. K. and Noel, J. P. (2012) The remarkable pliability and promiscuity of specialized metabolism. *Cold Spring Harbor Symp. Quant. Biol.* **77**, 309–320.
- Yokoyama, Y. and Esat, T. (2015) *Handbook of Sea-Level Research*. John Wiley & Sons, Chap. 7.

Yokoyama, Y., Suzuki, A., Siringan, F., Maeda, Y., Abe-Ouchi, A., Ohgaito, R., Kawahata, H. and Matsuzaki, H. (2011) Mid-Holocene palaeoceanography of the northern South China Sea using coupled fossil-modern coral and atmosphere-ocean GCM model. *Geophys. Res. Lett.* **38**, L00F03.

Zoccola, D., Tambutté, E., Kulhanek, E., Puverel, S., Scimeca, J. C., Allemand, D. and Tambutté, S. (2004) Molecular cloning and localization of a PMCA P-type calcium ATPase from

the coral *Stylophora pistilla*. *Biochim. Biophys. Acta Biomembr.* **1663**, 117–126.

SUPPLEMENTARY MATERIALS

URL (<http://www.terrapub.co.jp/journals/GJ/archives/data/52/MS511.pdf>)